# Earth's Internal Heat, Isotopes & Radioactive Decay



### **EARTH'S INTERNAL HEAT ENGINE AND CONVECTION**

The diagram below shows a cross section of Earth to a depth of approximately 700 km (not to scale). Arrows show motion of Earth's interior due to heat generated motion.



Wax cools down, contracts and sinks Wax heats up, expands and rises

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#### **CROSS-SECTION VIEW OF EARTH'S INTERNAL STRUCTURE**



Since its formation, Earth has been giving off heat through two processes:

- 1. Decay of radioactive materials
- 2. Residual heat from Earth's formation

The phenomenon of radioactivity relates to our story about the age of the Earth in two ways:

As radioactive elements decay in the Earth, they heat up the surrounding rocks.

(2) Radiometric Dating







# **ATOM**: the smallest particle that has the properties of an element.

Nucleus (99% of atom's mass):

uncharged <u>Neutrons</u> positively charged <u>Protons</u> Net charge of the nucleus is <u>POSITIVE</u>

**Electrons** in constant motion create a "cloud" like

a fan around the nucleus.

Charge of an electron is **NEGATIVE** 

### The Atom









 Isotopes are atoms that have the same # of protons, but a different # of neutrons.

*Number of Neutrons = Mass # - # of protons* 

 The difference in the number of neutrons cause isotopes to have <u>different mass numbers</u>

#### Mass Number = protons + neutrons

 Isotopes are unstable and fall apart releasing atomic particles – they <u>are radioactive</u>, release energy

## **Example of an Isotope**

• Example:

Carbon-12 (NORMAL) vs. Carbon-14 (ISOTOPE)



```
<sup>12</sup>C
Mass # = 12; Atomic # = 6
(6P, 6E, 6N)
How did we determine there were 6 neutrons?
```

```
<sup>14</sup>C
Mass # = 14; Atomic # = 6
(6P, 6E, 8N)
How did we determine there were 8 neutrons?
```



62 68 69 Lanthanide Pr Eu Tb Er Yb Ce Nd Pm Sm Gd Dy Ho Tm Serles 168.93 173.04 140.12 140.91 144.24 144.9) 150.36 151.97 157.25 158.93 162.5 164.93 167.26 91 93 94 95 96 97 98 99 102 Pa U Bk Th Np Pu Am Cf Es Fm Md No Cm (259.1) 232.04 231.04 258.1) 238.03 244.

Actinide Serles

> Alkali Metal Actinides Lanthanides Alkali Earth Non-metal Trans. Met. Halogen Noble Gas

Lu

174.97

Lr

Metal

### To estimate the age of a rock:

- D = amount of daughter product.
- P = amount of parent.

For a particular radioactive element in a rock, determine the present ratio = D/P.



 $\checkmark$  Rate of decay (from theory and measurement)

 $\checkmark$  Make assumptions about original ratios (from theory of geochemistry).



As time passes, the amount of parent decreases, and the amount of daughter product increases. This provides a way of estimating the amount of time since the "clock" got started (i.e., since the rock solidified).



Half Life = Number of years for 1/2 of the original number of atoms to decay from U to Pb

## Half-Lives of Radioactive Isotpes

Table 10.1 Radioactive isotopes frequently usedin radiometric dating.

Radioactive Parent	Stable Daughter Product	Currently Accepted Half-Life Values
Uranium-238	Lead-206	4.5 billion years
Uranium-235	Lead-207	713 million years
Thorium-232	Lead-208	14.1 billion years
Rubidium-87	Strontium-87	47.0 billion years
Potassium-40	Argon-40	1.3 billion years